

ICON – MPI – DWD

Icosahedral GCMs for climate research and operational weather forecasting

DWD: Günther Zängl, Daniel Reinert, et al.

MPI-M: Marco Giorgetta, Levy Silvers et al.

Former: Luca Bonaventura, Almut Gassmann, Hui Wan

Original motivation, ca. 2001

- MPI-M and DWD models are not satisfactory for future needs:
- MPI-M uses ECHAM/MPIOM as coupled climate model
 - ECHAM breaks tracer wind consistency due to a mix of spectral transform method for equations of motion and flux form semi-Lagrange transport scheme for tracers
 - ECHAM is hydrostatic → no option for cloud modelling
 - Scaling issues
- DWD uses global GME and regional COSMO for weather forecasting
 - Difficulties resulting from the usage of different dyn. equations, grids, numerics and parameterizations
 - No ocean model available for seasonal or decadal prediction

Goals and scope for new model

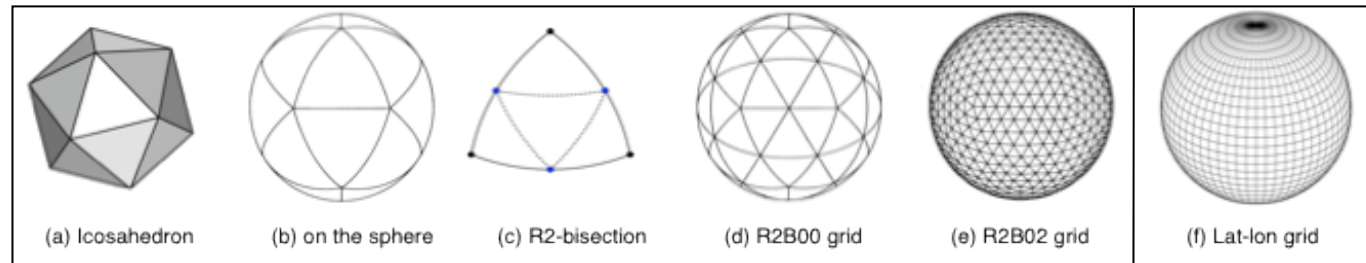
- Unified usage as climate model and NWP model to link daily verification of NWP to longer time scales
 - Atmospheric dynamical core must be fit for resolutions from ~1 to ~100 km and vertical extent to ~80 km
 - Tracer wind consistency
 - Regional refinement for simultaneous global and regional forecasting
 - Scale related physics packages (or scale aware physics?)
- High numerical efficiency and scalability:
 - Weak and strong scalability for NWP and climate
 - Usage on massively parallel machines (O(4+) cores)
- Framework for the development of atmosphere, land and ocean
 - Common infrastructure
 - Common grids



ICON characteristics

- Fully compressible equations of motion
- Mass and tracer mass conservation
- Icosahedral grids to avoid polar singularities and obtain a relatively uniform horizontal resolution
 - Option for triangle or hexagon/pentagon grids
- Static grid refinement by one-way or two-way grid nesting in 1 or several regions with individual refinement
 - Unstructured grid
 - Triangles allow simple refinement procedure
- Fast numerics that scales well and may be used with GPUs etc.
 - Small stencils
 - No global communication

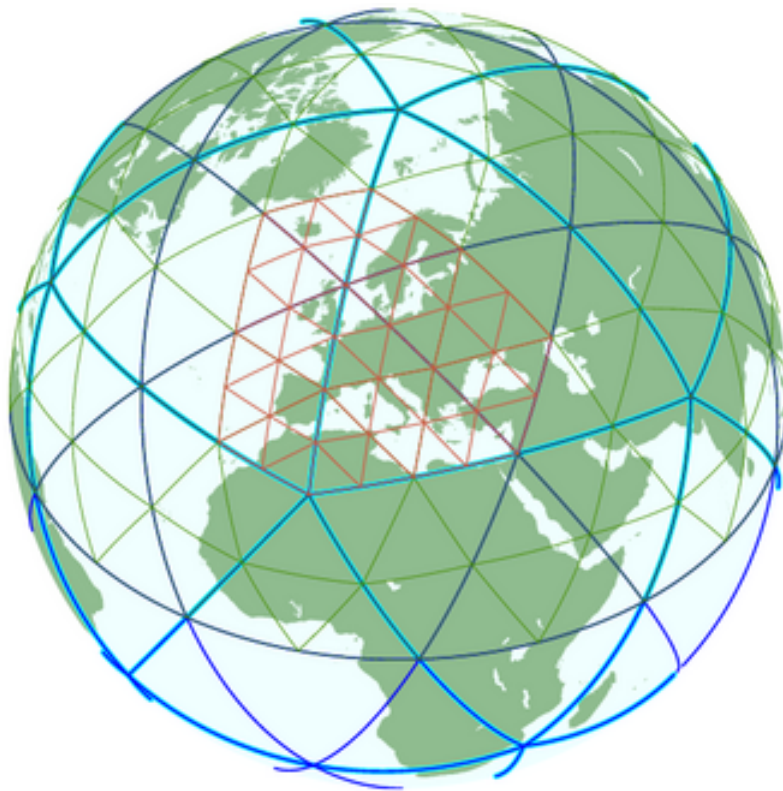
Grid construction



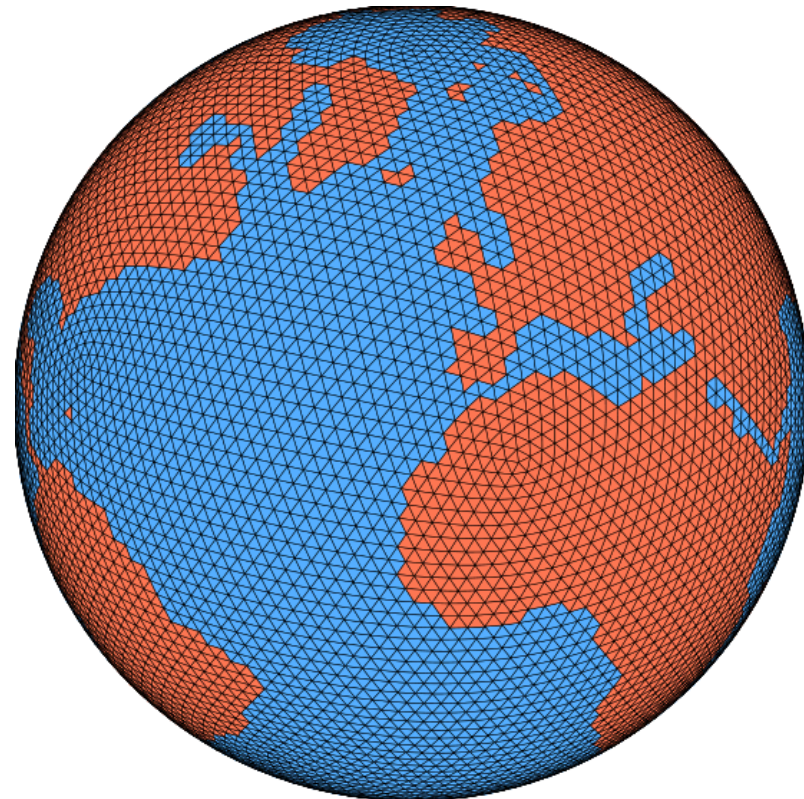
- (a) \rightarrow (b) Project the icosahedron into the sphere
- (c) and (d) Root refinement: Partition each edge into m sections of equal length and connect these points with great circle arcs parallel to the edges \rightarrow “ R_m ”
- (e) Further global refinements: Make bisection of edges and connect by great circle arcs, repeat n times \rightarrow “ B_n ”

Regional refinement: As for global, but limited to a region

Grid examples



- Light blue: spherical icosahedron (R1B0)
- Dark blue: global R2B0
- Green: Northern hemisphere R2B1
- Red: Europe R2B2

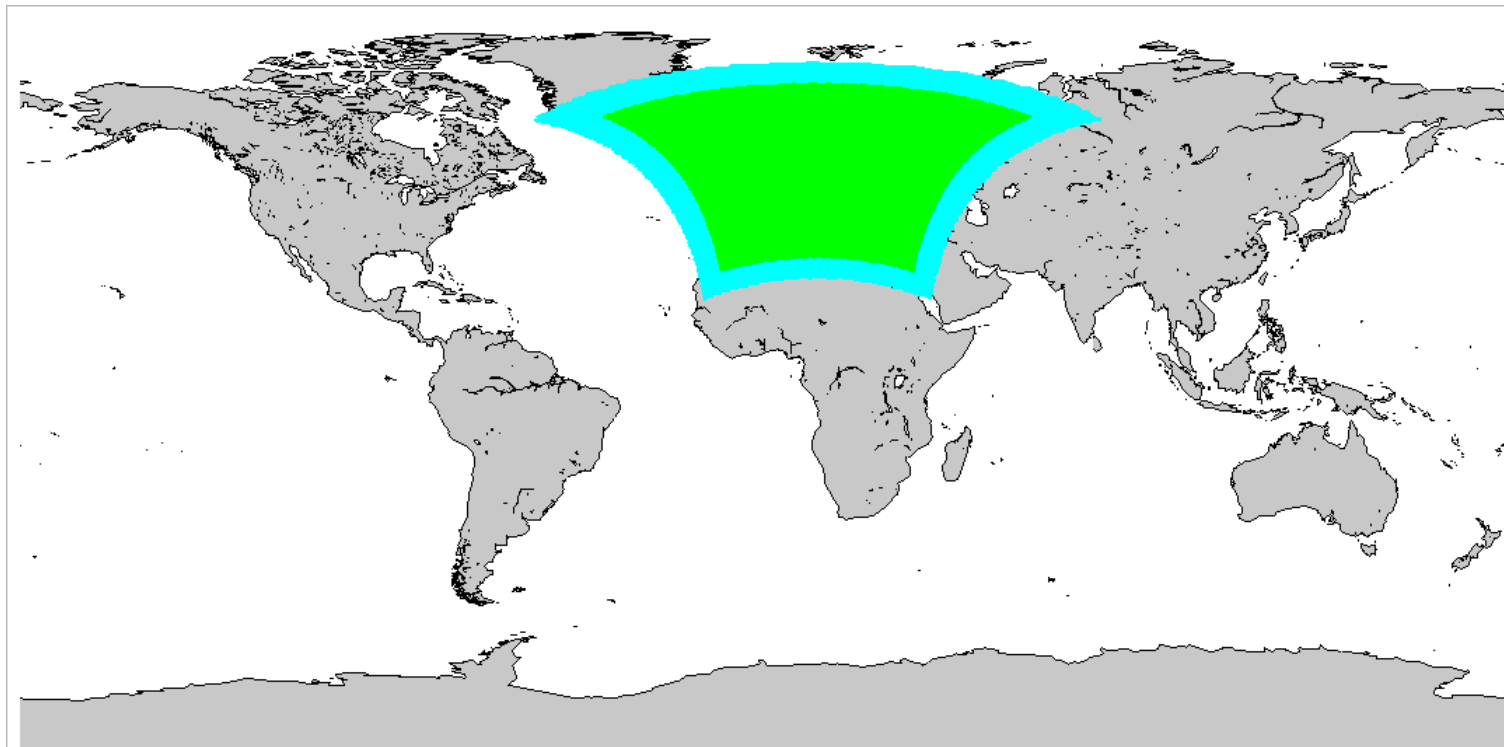


- Land sea mask at R2B4 (~140 km)
- Ocean cells have 2 wet edges
- Additional manipulation required for straits, passages etc.



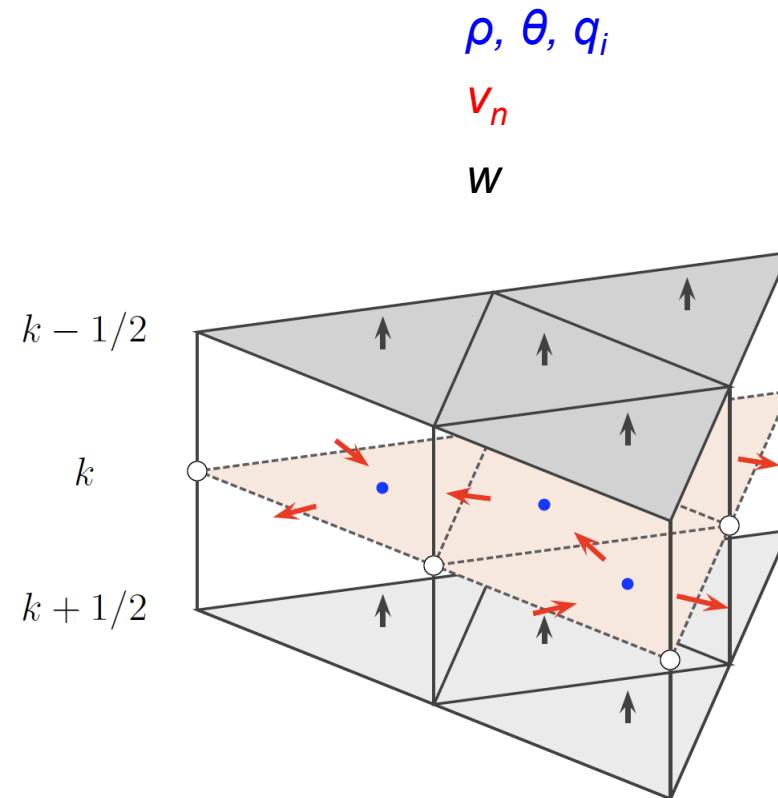
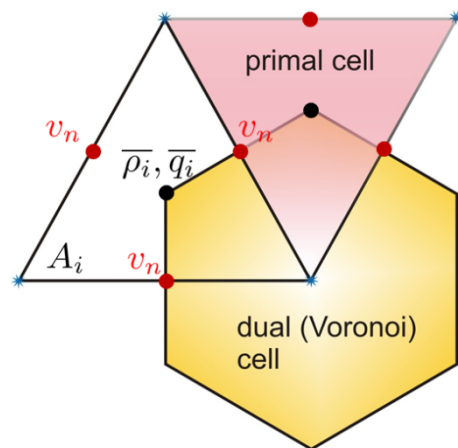
Prototype grid setup at DWD

- Global R2B7 (~20 km)
- 2 fold refinement over Europe (~5 km)



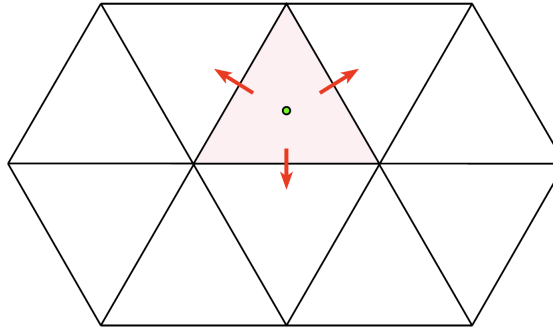
Staggering of variables

- C – grid staggering in horizontal and Lorenz grid in vertical:

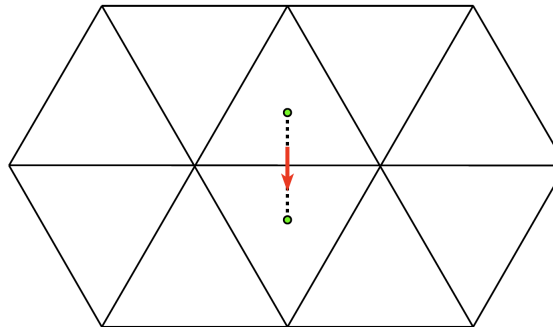


Basic operators on triangular grid

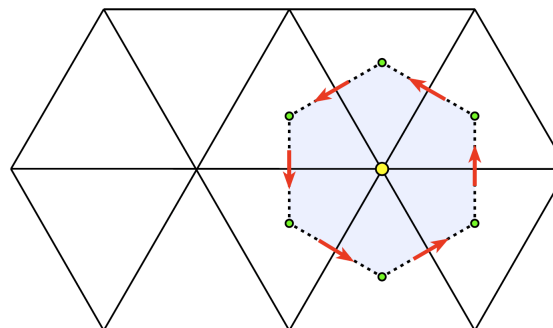
1. Divergence
Gauss's theorem
(problematic)



2. Gradient
finite difference

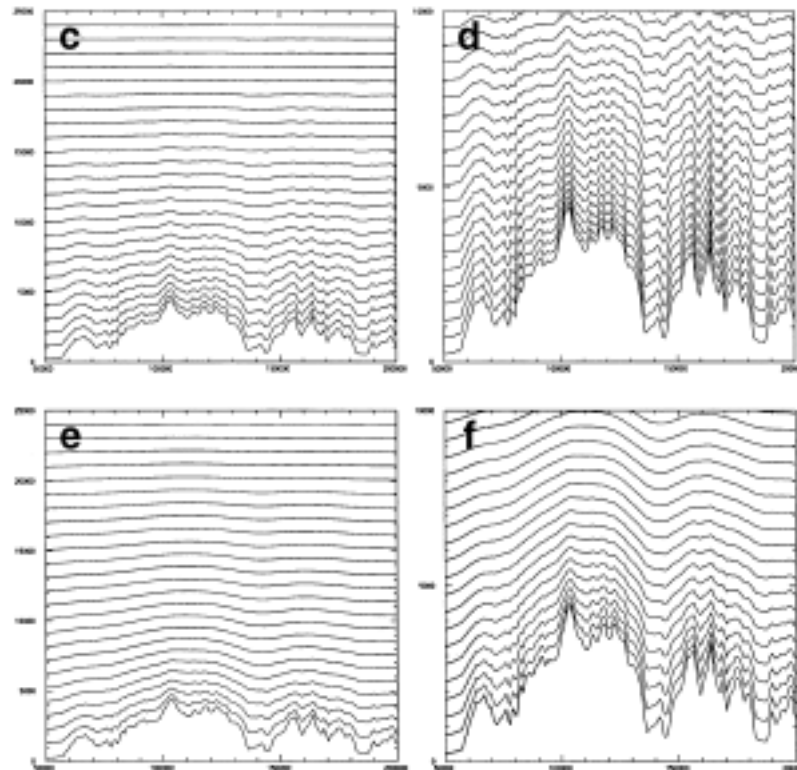


3. Curl
Stokes' theorem



Vertical grid

- Hybrid sigma height grid
 - Without or
 - or with smoothing with altitude (Schär et al., 2002)



(Fig. from Schär et al., 2002)



Fully compressible system of equations

(shallow atmosphere apr., spherical Earth, constant gravity)

$$\frac{\partial v_n}{\partial t} + (\zeta + f)v_t + \frac{\partial K}{\partial n} + w \frac{\partial v_n}{\partial z} = -c_{pd} \theta_v \frac{\partial \pi}{\partial n}$$

$$\frac{\partial w}{\partial t} + \nabla \cdot (\vec{v}_n w) - w \nabla \cdot \vec{v}_n + w \frac{\partial w}{\partial z} = -c_{pd} \theta_v \frac{\partial \pi}{\partial z} - g$$

$$\frac{\partial \rho}{\partial t} + \nabla \cdot (\vec{v} \rho) = 0$$

$$\frac{\partial \rho \theta_v}{\partial t} + \nabla \cdot (\vec{v} \rho \theta_v) = 0$$

v_n, w : normal/vertical velocity component

ρ : density

θ_v : Virtual potential temperature

K : horizontal kinetic energy

ζ : vertical vorticity component

π : Exner function

blue: independent prognostic variables



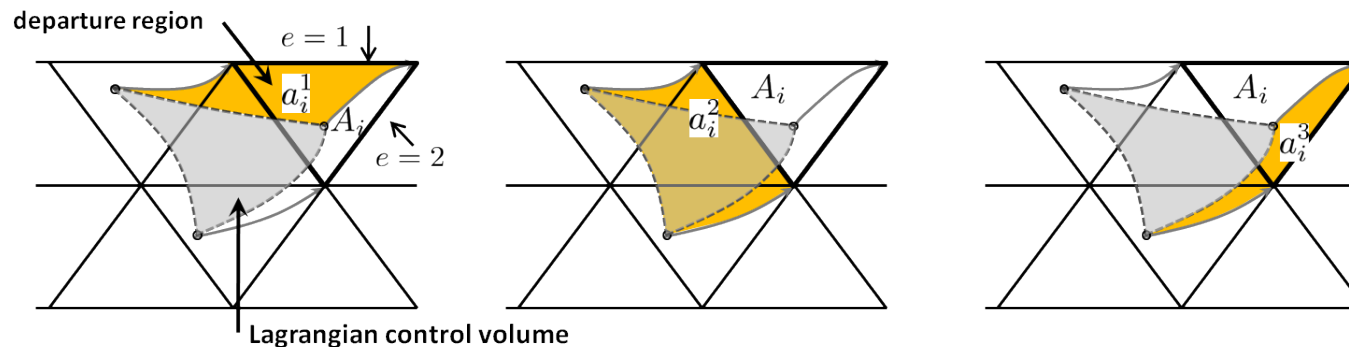
Numerical implementation

- 2D Lamb transformation for nonlinear momentum advection
- Flux form for continuity equation and thermodynamic equation; Miura 2nd-order upwind scheme (centered differences) for horizontal (vertical) flux reconstruction
- Finite-volume tracer advection scheme (Miura, 2007) with
 - 2nd-order and 3rd-order accuracy for horizontal advection;
 - 3rd-order piecewise parabolic method for vertical advection
 - monotonous and positive-definite flux limiters



Transport scheme (Miura, 2007)

- Compute flux of tracer mass over each edge from the approximated integrals of the tracer concentration over the departure regions



- Rhomboidal approximation of the departure region
- Linear, quadratic or cubic polynomial reconstruction
- Integrate polynomials with Gaussian quadrature (1 or 4 quadrature points).
- Simplified integration: apply polynomial of upwind cell, even if departure region overlaps with several neighboring grid cells
- Use dynamical core mass flux to obtain tracer and air mass consist

Time stepping

- Two time level predictor corrector scheme
- Horizontally explicit time stepping within the dynamical core at the time step needed for sound waves
 - Avoid complications of implicit methods and their risks for scaling properties
 - Efficiency gains of split-explicit time-stepping schemes (based on Leapfrog or Runge-Kutta) are less clear for global models extending to the mesosphere due to smaller ratio between sound speed and maximum wind speed
- Vertically implicit time stepping
- Longer time step for tracer advection and physics schemes

Process splitting

1. Dynamics with slow physics $n_{\text{now}} \rightarrow n_{\text{new}}^*$ (in m sub steps)
 - Radiation (reduced freq. for rad. transfer, red. resolution)
 - Convection
 - GWD
 - Divergence damping
2. Horizontal diffusion on v_n and θ : $n_{\text{new}}^* \rightarrow n_{\text{new}}^{*'}$
3. Tracer transport: $n_{\text{now}} \rightarrow n_{\text{new}}^{*'}$
 - Strang or Godunov splitting of horizontal and vertical advection
4. Fast physics, time split $n_{\text{new}}^{*'} \rightarrow n_{\text{new}}$
 1. Saturation adjustment
 2. Turbulent fluxes
 3. Cloud microphysics
 4. Saturation adjustment
 5. Surface/soil processes



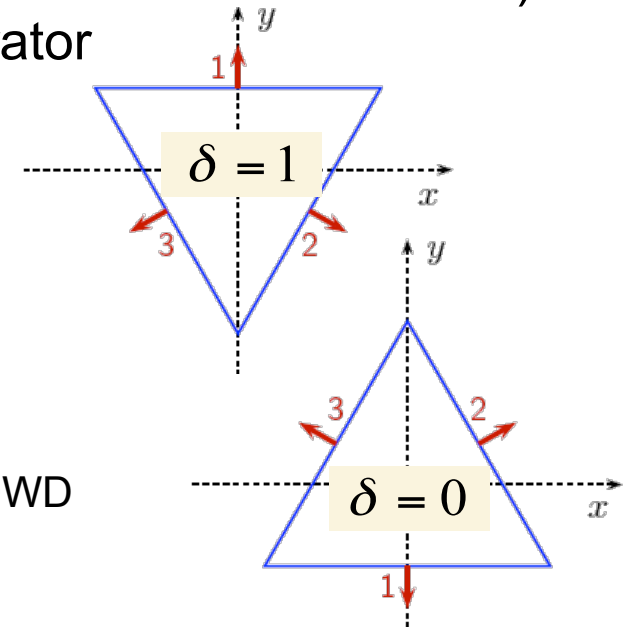
Problems

- Triangular C-grid has computational mode (cf. Dave Randall's talk) triggered by the discretized divergence operator

$$\text{div}(\mathbf{v})_i = (\nabla \cdot \mathbf{v})_o + \boxed{(-1)^\delta} l H(\mathbf{v})_o + \mathcal{O}(l^2)$$

- Options:

1. Suppress the computational mode \rightarrow ICON-MPI-DWD
 - Divergence damping
 - Filter normal wind used for divergence operator
2. Use hexagonal instead of triangular C-grid, \rightarrow ICON-IAP, Gassmann (2012)
3. Other grid/discretization \rightarrow cf. variety of DCMIP models



Summary

- The ICON-MPI-DWD model is developed for application over a wide range of scales in space and time
- Non-hydrostatic, fully compressible
- Mass and tracer mass conservation
- Numerical efficiency and high scalability
- Triangular C-grid problems suppressed numerically
- Model is in NWP test mode

END



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